

## Tectonophysics

**5150 Plate Tectonics**  
THE ORIGIN OF "HOTSPOT" TRACES: EVIDENCE FROM  
KALBAR, AUSTRALIA  
P. H. TILLEY, JR., (Geology Department, Louisiana  
State University, Baton Rouge, Louisiana, 70803)  
Interpretation of available tectonic data and  
geological evidence indicates a north-south  
transpressive pattern of extension of tectonic  
plates along a surveillance zone, beginning  
near 17°N. However, volcanic activity began  
about 15 Ma at 20°N and extending to the present,  
80 Ma over the length of the Mid-Atlantic Ridge  
and the Atlantic Ocean. The tectonic pattern of  
plate extension is consistent with that predicted by  
"hotspot" models. The trend of the trace is  
parallel with the Mid-Atlantic Ridge. The  
superior band is somewhat greater than 50 by 10  
km, relative to the Late Cretaceous and others (1977).  
The pattern of extension is consistent with the  
trend of the trace. The superior band is somewhat  
greater than 50 by 10 km, relative to the Late  
Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.  
The superior band is somewhat greater than 50 by  
10 km, relative to the Late Cretaceous and others  
(1977). The pattern of extension is consistent with  
the trend of the trace. The superior band is  
somewhat greater than 50 by 10 km, relative to the  
Late Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.

**5150 Plate Tectonics**  
THE ORIGIN OF "HOTSPOT" TRACES: EVIDENCE FROM  
KALBAR, AUSTRALIA  
P. H. TILLEY, JR., (Geology Department, Louisiana  
State University, Baton Rouge, Louisiana, 70803)  
Interpretation of available tectonic data and  
geological evidence indicates a north-south  
transpressive pattern of extension of tectonic  
plates along a surveillance zone, beginning  
near 17°N. However, volcanic activity began  
about 15 Ma at 20°N and extending to the present,  
80 Ma over the length of the Mid-Atlantic Ridge  
and the Atlantic Ocean. The tectonic pattern of  
plate extension is consistent with that predicted by  
"hotspot" models. The trend of the trace is  
parallel with the Mid-Atlantic Ridge. The  
superior band is somewhat greater than 50 by 10  
km, relative to the Late Cretaceous and others (1977).  
The pattern of extension is consistent with the  
trend of the trace. The superior band is somewhat  
greater than 50 by 10 km, relative to the Late  
Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.  
The superior band is somewhat greater than 50 by  
10 km, relative to the Late Cretaceous and others  
(1977). The pattern of extension is consistent with  
the trend of the trace. The superior band is  
somewhat greater than 50 by 10 km, relative to the  
Late Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.

**5150 Plate Tectonics**  
THE ORIGIN OF "HOTSPOT" TRACES: EVIDENCE FROM  
KALBAR, AUSTRALIA  
P. H. TILLEY, JR., (Geology Department, Louisiana  
State University, Baton Rouge, Louisiana, 70803)  
Interpretation of available tectonic data and  
geological evidence indicates a north-south  
transpressive pattern of extension of tectonic  
plates along a surveillance zone, beginning  
near 17°N. However, volcanic activity began  
about 15 Ma at 20°N and extending to the present,  
80 Ma over the length of the Mid-Atlantic Ridge  
and the Atlantic Ocean. The tectonic pattern of  
plate extension is consistent with that predicted by  
"hotspot" models. The trend of the trace is  
parallel with the Mid-Atlantic Ridge. The  
superior band is somewhat greater than 50 by 10  
km, relative to the Late Cretaceous and others (1977).  
The pattern of extension is consistent with the  
trend of the trace. The superior band is somewhat  
greater than 50 by 10 km, relative to the Late  
Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.  
The superior band is somewhat greater than 50 by  
10 km, relative to the Late Cretaceous and others  
(1977). The pattern of extension is consistent with  
the trend of the trace. The superior band is  
somewhat greater than 50 by 10 km, relative to the  
Late Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.

**5150 Plate Tectonics**  
THE ORIGIN OF "HOTSPOT" TRACES: EVIDENCE FROM  
KALBAR, AUSTRALIA  
P. H. TILLEY, JR., (Geology Department, Louisiana  
State University, Baton Rouge, Louisiana, 70803)  
Interpretation of available tectonic data and  
geological evidence indicates a north-south  
transpressive pattern of extension of tectonic  
plates along a surveillance zone, beginning  
near 17°N. However, volcanic activity began  
about 15 Ma at 20°N and extending to the present,  
80 Ma over the length of the Mid-Atlantic Ridge  
and the Atlantic Ocean. The tectonic pattern of  
plate extension is consistent with that predicted by  
"hotspot" models. The trend of the trace is  
parallel with the Mid-Atlantic Ridge. The  
superior band is somewhat greater than 50 by 10  
km, relative to the Late Cretaceous and others (1977).  
The pattern of extension is consistent with the  
trend of the trace. The superior band is somewhat  
greater than 50 by 10 km, relative to the Late  
Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.  
The superior band is somewhat greater than 50 by  
10 km, relative to the Late Cretaceous and others  
(1977). The pattern of extension is consistent with  
the trend of the trace. The superior band is  
somewhat greater than 50 by 10 km, relative to the  
Late Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.

**5150 Plate Tectonics**  
THE ORIGIN OF "HOTSPOT" TRACES: EVIDENCE FROM  
KALBAR, AUSTRALIA  
P. H. TILLEY, JR., (Geology Department, Louisiana  
State University, Baton Rouge, Louisiana, 70803)  
Interpretation of available tectonic data and  
geological evidence indicates a north-south  
transpressive pattern of extension of tectonic  
plates along a surveillance zone, beginning  
near 17°N. However, volcanic activity began  
about 15 Ma at 20°N and extending to the present,  
80 Ma over the length of the Mid-Atlantic Ridge  
and the Atlantic Ocean. The tectonic pattern of  
plate extension is consistent with that predicted by  
"hotspot" models. The trend of the trace is  
parallel with the Mid-Atlantic Ridge. The  
superior band is somewhat greater than 50 by 10  
km, relative to the Late Cretaceous and others (1977).  
The pattern of extension is consistent with the  
trend of the trace. The superior band is somewhat  
greater than 50 by 10 km, relative to the Late  
Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.  
The superior band is somewhat greater than 50 by  
10 km, relative to the Late Cretaceous and others  
(1977). The pattern of extension is consistent with  
the trend of the trace. The superior band is  
somewhat greater than 50 by 10 km, relative to the  
Late Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.

## Volcanology

**5150 Plate Tectonics**  
THE ORIGIN OF "HOTSPOT" TRACES: EVIDENCE FROM  
KALBAR, AUSTRALIA  
P. H. TILLEY, JR., (Geology Department, Louisiana  
State University, Baton Rouge, Louisiana, 70803)  
Interpretation of available tectonic data and  
geological evidence indicates a north-south  
transpressive pattern of extension of tectonic  
plates along a surveillance zone, beginning  
near 17°N. However, volcanic activity began  
about 15 Ma at 20°N and extending to the present,  
80 Ma over the length of the Mid-Atlantic Ridge  
and the Atlantic Ocean. The tectonic pattern of  
plate extension is consistent with that predicted by  
"hotspot" models. The trend of the trace is  
parallel with the Mid-Atlantic Ridge. The  
superior band is somewhat greater than 50 by 10  
km, relative to the Late Cretaceous and others (1977).  
The pattern of extension is consistent with the  
trend of the trace. The superior band is somewhat  
greater than 50 by 10 km, relative to the Late  
Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.  
The superior band is somewhat greater than 50 by  
10 km, relative to the Late Cretaceous and others  
(1977). The pattern of extension is consistent with  
the trend of the trace. The superior band is  
somewhat greater than 50 by 10 km, relative to the  
Late Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.

## General or Miscellaneous

**5150 Plate Tectonics**  
THE ORIGIN OF "HOTSPOT" TRACES: EVIDENCE FROM  
KALBAR, AUSTRALIA  
P. H. TILLEY, JR., (Geology Department, Louisiana  
State University, Baton Rouge, Louisiana, 70803)  
Interpretation of available tectonic data and  
geological evidence indicates a north-south  
transpressive pattern of extension of tectonic  
plates along a surveillance zone, beginning  
near 17°N. However, volcanic activity began  
about 15 Ma at 20°N and extending to the present,  
80 Ma over the length of the Mid-Atlantic Ridge  
and the Atlantic Ocean. The tectonic pattern of  
plate extension is consistent with that predicted by  
"hotspot" models. The trend of the trace is  
parallel with the Mid-Atlantic Ridge. The  
superior band is somewhat greater than 50 by 10  
km, relative to the Late Cretaceous and others (1977).  
The pattern of extension is consistent with the  
trend of the trace. The superior band is somewhat  
greater than 50 by 10 km, relative to the Late  
Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.  
The superior band is somewhat greater than 50 by  
10 km, relative to the Late Cretaceous and others  
(1977). The pattern of extension is consistent with  
the trend of the trace. The superior band is  
somewhat greater than 50 by 10 km, relative to the  
Late Cretaceous and others (1977). The pattern of  
extension is consistent with the trend of the trace.

## The Challenge of Climate to Man

Alan D. Hecht

Climate Dynamics Program  
National Science Foundation  
Washington, D.C.

How vulnerable is the United States and world food supply to a serious drought today? Will the burning of fossil fuel and the subsequent release of CO<sub>2</sub> to the atmosphere alter global climate? Is society today sufficiently resilient to respond to major climatic changes? Is there a coming ice age?

### The Climate Challenge

Around 900 A.D. a group of small villages was established in northwest Iowa by Indians of what we now call the Mill Creek culture. Around 1400 A.D., after many prosperous years, the villages were abandoned. In the summer of 1863, archaeological and geological excavations of several sites of the Mill Creek culture began. While three major sites were excavated, one known as Phipps site provided the clearest historical record of civilization in the area. To reconstruct both the habits of the Mill Creek people and the environment in which they lived, scientists have studied a wide assortment of remains preserved in the strata of northwestern Iowa. Unnoticed without the aid of a microscope are the remains of pollen grains blown into the area from surrounding trees. The pollen preserved in the strata can be read as an historic log of changes in the vegetation and climate surrounding the Mill Creek area. The village was occupied about 900 A.D. on the flood plain of Mill Creek. The pollen evidence shows that during the 10th and 11th centuries, the Indians lived in a region with tall-grass prairie on the uplands and woods on the valley terraces and valley floors. This vegetation is not very different from today's it one substitutes "cornfield" for "prairie." From evidence given by fossil bones found in the strata, it seems that deer and elk were abundant and were hunted by these Indians. The Indian meat diet appears to have been dominated by these animals, supplemented only occasionally by bison. Maize was the main agricultural product.

Toward the end of the 12th century major environmental changes occurred at Mill Creek. The influx of oak pollen began to decline rapidly, while populus (probably cottonwood) rose rapidly. The proportion of bison meat eaten by the Indians rose abruptly at this time. Within a few decades in the 12th century the vegetation in the entire area changed from tall-grass prairie on the uplands and forest in the larger valleys to steppe-like vegetation and essentially only phreatophytes along the streams in all but the major

valleys. From radiocarbon-dated samples of charcoal and from the accumulation rate of sediment at the site, the rate of pollen changes in this area can be estimated. The decline of oak pollen from its maximum to minimum occurred in less than a century. The rapid rise of grass pollen took about 45 years; the rise of phreatophytes about 15 years.

The interpretation of the changes in pollen preserved in the Mill Creek area and the changes in feeding habits of the Mill Creek Indians suggest the beginnings of a long-term drought. In perhaps one or two generations (45 years) the tall-grass prairie was replaced by short grass. The few cottonwoods and willows along the stream banks were the only remnants of the forest that once filled the valleys. The deer, a woodland browser, disappeared, and two thirds of the meat eaten by the Mill Creek people came from bison, a short-grazing animal. Further west of the Mill Creek area, other archaeological evidence indicates that the farming villages disappeared entirely.

The Mill Creek site has been extensively studied by Reid Bryson and his colleagues at the Institute for Environmental Studies of the University of Wisconsin [Bryson and Baser, 1968; Bryson et al., 1970]. Their documentation of the drought conditions in this area during the 12th to the 14th century is relevant to society today since this area is now a major spring wheat, maize, and soybean region for the United States. The drought at Mill Creek forced the abandonment of a corn-farming community which had lasted for 500 years.

Today, centuries later, in a highly developed technological society, we still face problems similar to those of the Mill Creek Indians, although we possess much greater powers of hindsight and foresight in the matter of climate variability and change. There is growing apprehension, for example, that man-made increases in atmospheric CO<sub>2</sub> are contributing to a global climate warming on a scale yet to be experienced in historical times. There is some scientific evidence to suggest that such a change could spell a gradual warming and drying of the environment once occupied by the Mill Creek Indians and now the center of U.S. agricultural production. In the more immediate future, there is renewed concern over the possibility of a recurrence of a severe drought, an event which threatens sudden disruption to an increasingly global food system.

Problems of both long-term climate change and short-term variability—of CO<sub>2</sub> and drought in particular—are explored at greater length in this essay.

### Drought in the Great Plains

Man and drought have been at odds since the dawn of civilization. In the continental U.S., droughts have been known, according to historical documents, since the early 1800's [Ludlum, 1971].

To a first approximation, droughts have occurred in the midwestern U.S. and the Great Plains in particular, roughly every 20 years, although their distribution and intensity have been quite different for each drought period. For example, the droughts in 1910, 1911, 1913, and 1917 were short, severe, and spatially limited, as was the drought in the 1960's. The drought of the 1930's, however, was widespread and persistent. No drought since has equaled the intensity, areal extent, and persistence of the drought of the 1930's.

The severity and duration of aridity in the area can be related to moisture balance by a meteorological index developed by Palmer [1965]. The Palmer Drought Severity Index (PDSI) is based on an empirical water balance approach. The normal amount of precipitation received in an area is dependent on the average climate and the meteorological conditions of the area both during and preceding the month or period in question. The Palmer index computes the required precipitation for any area. The difference between the actual and computed precipitation is a measure of the deviation of the amount of moisture from the long-term mean. The index is structured to correspond to a wide range of moisture conditions, as shown in Table 1.

TABLE 1. Drought Severity Index (PDSI)		
Palmer Index		Degree of Drought
-4.0 <	≤ -4.0	Extremely dry
-3.0 <	≤ -3.0	Severely dry
-2.0 <	≤ -2.0	Moderately dry
-1.0 <	≤ -1.0	Mildly dry
-0.5 <	≤ -0.5	Near Normal
+0.5 <	≤ +0.5	Mildly wet
+1.0 <	≤ +1.0	Moderately wet
+2.0 <	≤ +2.0	Severely wet
+3.0 <	≤ +3.0	Extremely dry

Classification of moisture conditions, based on a scheme developed by the meteorologist, W. C. Palmer. Index refers to meteorological rather than soil conditions.

The PDSI is one of several drought indices calculated by Department of Commerce/NOAA for the entire U.S. The index can be read as a measure of local areal moisture that is relative to the long-term mean. The formula for making the calculation also includes a "memory" term for conditions during previous months. An important property of the PDSI is that the same number in different locations means roughly the same relative degree of drought.

Figure 1, for example, shows the reconstructed PDSI values for 64 climatic divisions in the Great Plains for the period 1931 to 1977 [Warwick, 1980]. These data show that the drought of the 1930's has not been duplicated, in both intensity and duration, by subsequent droughts. The 1950's drought matched that of the 1930's in severity but was limited to certain portions of the central and southern Plains. Isolated drought occurred in the 1960's and 1970's. The fact that there is a well-known (but poorly understood) sunspot cycle of nearly 22 years and drought occur-

## PATTERNS OF GREAT PLAINS DROUGHTS

Based on Palmer Index Values averaged over four months, May through August, by climatic division.

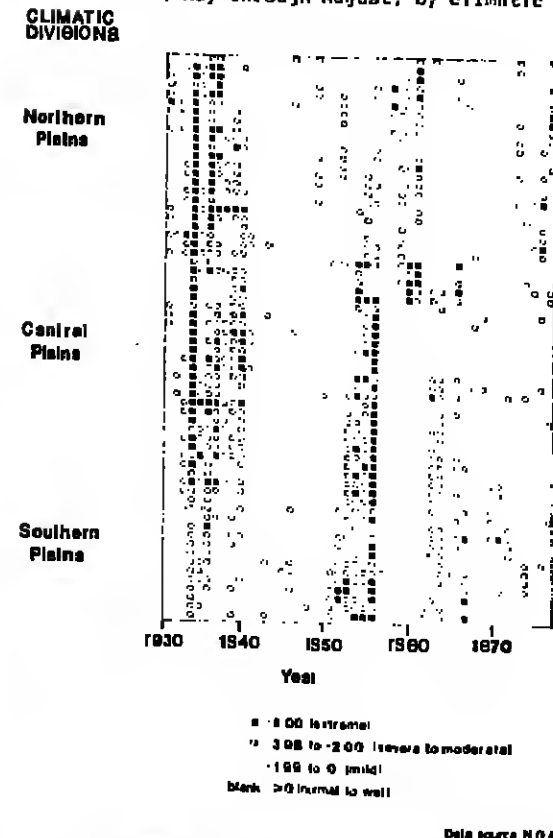


Fig. 1. Patterns of Great Plains droughts by geographic division, based on PDSI values averaged from May through August (From Warwick [1980]; reproduced with permission of the author.)

rence in the Great Plains of nearly every 20 years is often the basis for postulating links between these activities. Mitchell et al. [1979] have provided some empirical evidence for coincidence in frequency between sunspot activity and drought intensity. These authors built their analysis on the established relationship between variation in the width of tree rings and climate [Fritts, 1976]. In their study the variations in tree ring width for the western U.S. was correlated with calculated values for PDSI. An equation relating the two variables was derived and used to determine PDSI for times before meteorological records. In the end, PDSI were determined for approximately 40 localities west of the Mississippi River for the period from 1700 to the present. Areas where the PDSI were of equal values (-1, -3, -2, and -1) then were grouped for each year to produce a Drought Areas Index (DAI). This index was then analyzed by spectral techniques. The dominant frequency identified in these series was 22 years. Thus both sunspot activity and drought frequencies in the western U.S. have the same frequency. Mitchell et al.'s detailed analysis and conclusion provide an excellent perspective on what this coincidence means.

From the viewpoint of solar physics and solar terrestrial mechanisms of potential relevance to climate, our results would clearly seem to imply a role of solar magnetic activity in giving rise to widespread drought in the western U.S. This role may be either direct or indirect. It is our impression that the solar control of drought is not to be construed as a prime mover of drought or of climatic aberrations that result in drought. Rather, we prefer to think that the solar control is in the nature of a modulating mechanism, that alternatively favors or discourages the spread of drought at times when terrestrial climatic developments unrelated to solar events are primed to erupt into a drought situation.

These results provide no justification for using solar variability as a reliable basis for climate or drought prediction. Our data make it abundantly clear that a wet year can arrive at a time when the Sun "says" it should be a drought year, and that a major drought can develop when the sun "says" there should be no drought.

### Drought and International Politics

The lessons of Mill Creek and the historic records of drought in the Midwest underscore the recurrent nature of drought and its impact on society. From the time of the first subsistence in this region, in the late 1800's, to today, the Great Plains has grown in importance as a major food-producing area. It accounts for 12%-15% of the world's total wheat production and 61%-65% of the nation's wheat. The U.S. also provides 40%-45% of the world's total wheat trade. This blessing from the land is the product of sophisticated technology and a generally favorable climate over the last 100 years.

The question of how much each of these variables (technology and weather/climate) affect crop yield is controversial and unresolved. It is an extremely important question, however, since it affects the types of management strategies used in agricultural decisions. One school of thought maintains that the sensitivities of crop yields to drought have

Figures are based on 1975-1978 in Agricultural Statistics, 1979 USDA.

The National Climate Program Act: Hearings before the Subcommittee on the Environment and the Atmosphere of the Committee on Science and Technology (94th Congress, 2nd Session).

## EOS

TRANSACTIONS, AMERICAN GEOPHYSICAL UNION

The Weekly Newspaper of Geophysics

Send double-spaced manuscripts (four copies) to Eos, AGU, 2000 Florida Avenue, N.W., Washington, D.C. 20008, or send them directly to one of the associate editors with a copy to the above address.

**Editor:** A. F. Spilhaus, Jr.; **Associate Editors:** Claudia J. Allegre, Peter M. Bell, Kevin C. Burke, Arnold L. Gordon, Kristina Karam, Gerard Lachapelle, Christopher T. Russell, Richard A. Seifert, Sean C. Solomon, Carl Kisslinger; **News Writers:** Barbara T. Richmond; **Editor's Assistants:** Sandra R. Marks; **Eos Production Staff:** Patricia Bangert, Margaret W. Connolly, Eric Garfield, James Hobbethwaite, Dee Sung Kim, Michael Schwartz.

**Officers of the Union**  
J. Tuzo Wilson, President; James A. Van Allen, President-Elect; Leslie H. Meredith, General Secretary; Carl Kisslinger, Foreign Secretary; A. F. Spilhaus, Jr., Executive Director; Waldo E. Smith, Executive Director Emeritus.

Advertising that meets AGU standards is accepted. Contact Robin E. Little, advertising coordinator, 202-482-8903.

Eos, Transactions, American Geophysical Union (ISSN 0096-3941) is published weekly by the American Geophysical Union from 2000 Florida Avenue, N.W., Washington, D.C. 20008. Subscription available on request. This issue \$5.00. Second-class postage paid at Washington, D.C., and at additional mailing offices.

Copyright 1981 by the American Geophysical Union. Material published in this issue may be photocopied by individual scientists for research or classroom use. Permission is also granted to use short quotes and figures and tables for publication in scientific books and journals. For permission for any other uses, contact AGU Publications Office, 2000 Florida Avenue, N.W., Washington, D.C. 20008.

Views expressed in this publication are those of the authors only and do not reflect official positions of the American Geophysical Union unless expressly stated.

**Cover:** Schematic illustration of the components of the climate system. Dark arrows are examples of external processes; open arrows are examples of internal processes in climatic changes. (From a U.S. Climate Program Plan, NOAA; Department of Commerce.)



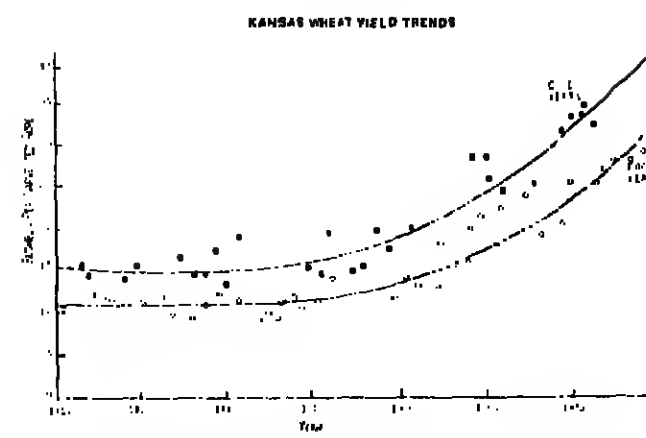


Fig. 2. Trends in wheat yield in Kansas. (From Warrick [1980]; reproduced with permission of the author.)

been reduced (over the past 30 years) because of advances in technology (see, for example, Newman, 1978). A second school of thought suggests that crop yields are sensitive to sharp declines from drought, even given new technological advances (as, for example, McQuigg et al., 1973). A major barrier to resolving this question is the stubborn problem of separating weather and technological factors in agricultural production. It is probably unrealistic to expect a solution to the problem when comparisons are made at the level of large geographic areas that cross climatic and/or geologic zones. Warrick (1980) suggests a different approach to the problem that focuses on states, crop-reporting districts (which coincide with state climatic divisions for the Great Plains), or counties. His analysis of wheat yield in Kansas over the period 1930 to the present suggests a sensitivity to drought conditions. Figure 2, for example, shows yield trends in Kansas separated by good weather and bad weather years. If technological improvements in yield type or in management have occurred over this period, it might be expected that the two curves for good and bad weather years would in time converge. In other words, if the agricultural system had become better in dealing with drought conditions, the relative difference between expected good yield and expected poor yields would decrease over time, but they do not. In terms of absolute bushels the curves in Figure 2 actually diverge. This is a warning, at least for wheat production in Kansas, that agriculture has not completely engineered climate out of the picture of crop production.

As the Midwest is an important food source, a danger lies ahead in not knowing how resistant the area is to a recurrence of a 1930's-type drought. Richard Warrick at the National Center for Atmospheric Research and his collaborators at Clark University are addressing one of the more important questions of the time, namely what would be the global impact if such a drought occurred in the Great Plains. Warrick's important preliminary findings on linking climate yield and global food trade models suggest that a recurrence of a 1930's drought in today's world might induce famines in grain-import-dependent regions that would exceed, in total deaths worldwide, any similar catastrophe since the 1930's. Further model-linking analyses are being performed to explore this question in greater detail (Warrick and Kates, 1981).

The relationship between climate and society today is far more complex than during Mill Creek times. Complex management decisions and political, economic, and social patterns today can serve to increase or lessen the environmental impact of climate change. The 1968-1973 drought in semiarid West Africa (Sahel) is a case in point.

Societal scientists have documented that political, economic, and agricultural factors were partly responsible for the magnitude of the crisis in the Sahel (Glanz, 1977). Man may have helped create or intensify the drought by destruction of vegetation which, in turn, increased surface albedo

and thereby decreased rainfall (Chamey, 1975). This process can turn marginal lands, such as the Sahel, into deserts.

The lesson of Mill Creek, in a broad sense, is to underline the relationship between climate and man. Climate can be thought of as a natural resource, a concept originally developed by Landsberg (1948). How society responds to climatic fluctuation, how it manages its resources in light of climatic change, and how it may alter global climate by its own activities may well be measured, by the year 2000, in economic terms, population increase, and perhaps world famine. National and international programs are now being developed to understand better the role of climatic processes in shaping the world's economy and welfare.

### Policy implications

The decade of the 1970's was characterized by sufficiently adverse social and weather conditions in many parts of the world as to suggest to many in policy, management, and government positions that more attention should be paid to understanding the impacts of climate on society. A 1974 report from a committee of the existing Domestic Council (A United States Climate Program) said:

The food and energy crisis is being sharply intensified throughout the world by the natural fluctuation of regional climate. Longer-term changes in climate, whether naturally occurring or resulting from man's activities or both, may be leading to new global climate regimes with widespread effects on food production, energy consumption, and water resources. These circumstances have created an urgent need for a program that can offer hope of knowing and anticipating the effects of climate fluctuations and changes here (U.S.) and around the world. A U.S. Climate Program is proposed which will enable the U.S. Government to meet this need.

Between 1974 and 1977, while numerous government committees and the National Academy of Sciences developed various aspects of an integrated U.S. climate program, the U.S. Congress began considering legislation for the initiation of a national climate program. On May 18, 1978, the House Subcommittee on the Environment and the Atmosphere (of the Committee on Science and Technology) met under the chairmanship of Congressman George Brown (D., Calif.) for the first of 8 days of hearings on the subject of climate and related research.

Congressman Brown's opening remarks reiterated the theme that the 1970's were turbulent social and climatic times.

I am sure that events in recent years have made us all aware of the impacts of climate on mankind. Perhaps the most memorable event was the drought in Russia in 1972, which led to the infamous grain sale. Along with the concurrent failure of the Peruvian anchovy fishery due to a changing ocean current, this was one of the major causes of the spectacular inflation in food prices during 1972 and 1973. More recently, we have seen the effects of a disastrous drought in the Sahel, failure in the Indian Monsoon, and closer to home, a drought in the northern part of California which is badly affecting this year's crops.

Despite the above perception, there is actually no evidence that climate everywhere is becoming more variable. Chico and Sellers (1979), for example, have examined the variability of mean monthly temperatures in the United States since 1898. Their results show that the interannual variability was greatest in the decade centered on 1930, and it has decreased steadily to a minimum in the decade centered around 1970. This trend in variability is almost completely explained by changes in variability during the

winter months of December to February. The great change in variability for the U.S. occurred in the Midwest. Even if variability has not changed significantly over the past decades, the effects of climate variability have been felt on society through economic and social hardship.

For example, the economic impact of the abnormally cold winter of 1976-1977 in the eastern half of the United States in agricultural losses alone was approximately \$2 billion. [Source: State Government News, April 1977, published by Council of State Governments.]

### Estimated Crop and Capital Investment Losses During Winter 1976-1977

Arkansas—\$39 million total losses, including soybeans and hay. Future pasture production could require lengthy recovery. Georgia—Cattle producers hurt, pastures diminished. Indiana—\$5.87 million loss to livestock, which will be difficult to recover, and another \$10.5 million agricultural loss, including milk dumped. Kentucky—\$108 million total losses, much of it to livestock and to increased feed and labor cost for livestock. Louisiana—\$60 million in cattle and crop losses, with long-range figure much higher. Many cattlemen sold foundation stock. Much of sugar cane and citrus crop lost. Massachusetts—Problems with transportation of produce and feed.

Maryland—Agricultural losses of \$25 million, including livestock, broilers, wheat, and tobacco. Seafood industry lost 40-50 days of harvest time in Chesapeake Bay that will have a long-range effect on oyster and blue crab industry. Michigan—\$158,000 in milk dumped because of snow-blocked roads. Mississippi—Excess of \$100 million in losses, chiefly in cattle industry. Following months of unprofitable cattle operations, the winter caused a severe strain on the ability of cattlemen to recover. Stress on breeding herds will be felt a long time. New York—At least \$3.5 million in agricultural damage, some \$65 million worth of milk dumped, and \$750,000 dairy cattle lost because of barn collapses. Ohio—Total loss of \$15.2 million, including 93,000 livestock. Pennsylvania—Milk dumped; peach, winter wheat, barley, and all other crops affected; pigs sold at loss; increased feed costs and barn cave-ins. South Carolina—Total loss of \$41.2 million in feed and cattle. Livestock producers need 93,363 tons of hay and 1.3 million bushels of grain to maintain herds. Request for federal aid denied. Tennessee—Up to \$10 million in losses. Virginia—Total reduction in value of farm production of \$150-\$160 million, including 1.2 acres of hay and pasture affected and crop, nursery, livestock, and capital investment losses. Potential farm income reduced by 11%.

Total monetary losses \$2,356.6 million.

The 1980 summer drought in the Midwest and south central United States has also had significant economic and health effects. While the full impact of this drought is not yet known, the heat and lack of moisture has reduced crop yield significantly below previous year yields, and total estimated crop losses are over \$1 billion [State Government News, Aug. 1980].

While the 94th Congress considered the need for a national climate program, no legislation was successfully developed. One year later, in June 1978, additional hearings were held by the Committee on Commerce, Science and Transportation of the U.S. Senate.

The 95th Congress eventually passed a National Climate Program Act (P.L. 95-387), which was signed by President Carter on September 17, 1978. The act is designed to establish a comprehensive and coordinated national climate program. The act is a realization that the effects of climate have important social, economic, and political consequences and that should be given consideration in policy and resource planning. A 5-year plan to implement specific goals of this national plan has recently been prepared by the National Climate Program Office.

While, since 1974, the U.S. has promoted the concept of a national climate program, similar developments have been underway in Europe. In February 1979, the World Meteorological Organization (WMO) convened the first 'World Climate Conference,' as a beginning of a World Climate Program (WCP). The WCP, which became effective in January 1980, will now be the focus for large-scale international programs in climate research and service.

Additionally, the United Nations, through its environmental program (UNEP) is taking a lead role in promoting programs to study the impact of climate on society. One major issue on UNEP's agenda is the impact of CO<sub>2</sub> on climate and the resulting impacts on society.

### The Climate System: Recognizing Signal From Noise

The effect of CO<sub>2</sub> or any other anthropogenic influence on climate must be distinguished above the natural background of climate variability. Climate varies on all time scales, only a sampling of which is discussed below.

The earth's climate is characterized by its constant state of flux. It is the product of the interactions of the atmosphere, oceans and cryosphere, and the earth's surface. The cover figure shows a simple representation of the processes operating within what may be called the climate system, and processes operating on the climate system. Radiation from the sun provides the fundamental energy that drives this system. The variation of chemical constituents of the atmosphere, such as CO<sub>2</sub>, aerosol, dust, etc. to change the amount of radiation incident on the earth's surface. The radiation, once received at the earth's surface, drives the atmospheric circulation, which in turn drives the oceanic circulation. The oceanic circulation is closely linked to the circulation of the atmosphere. Together, the interaction of the atmosphere and oceans are influenced by the extent and thickness of the ice covering the land and sea.

Although weather and climate are sometimes used interchangeably, there are important distinctions between them.

Weather is the state of the atmosphere (described as completely as possible with present observing capabilities) at one point in time. Weather prediction attempts to forecast a new condition of the atmosphere—given an initial atmospheric state—by the application of fundamental laws of atmospheric motion.

Climate results from an ensemble of weather events for a season, year, or longer period. A climate state is usually defined in terms of average conditions as well as some measure of the variability within the time period under consideration.

Although the same physical laws apply to both weather and climate, climate prediction is complicated by the need to consider complex interactions (as well as changes within) all components of the earth's climate system. For example, while it may not be necessary to consider the small changes in ocean temperature or circulation from one day to another for a successful weather forecast, such changes become important when predicting atmospheric changes from one season to another. Similarly the changes in the geometry of the earth's orbit occur on a time scale that is important for deducing climate changes over thousands of years, but they are of no consequence to weather forecasting.

### Climate Variability: The Last 100 Years

A summary of the major features of climate variability on several different time scales is shown in Figure 3. Northern Hemisphere average temperatures for the past 100 years show a general trend of increasing temperatures from the 1880's to the 1940's, and declining temperatures thereafter. These temperature changes are on the order of tenths of a degree, although the change from 1880 to 1940 is a change of nearly one full degree centigrade.

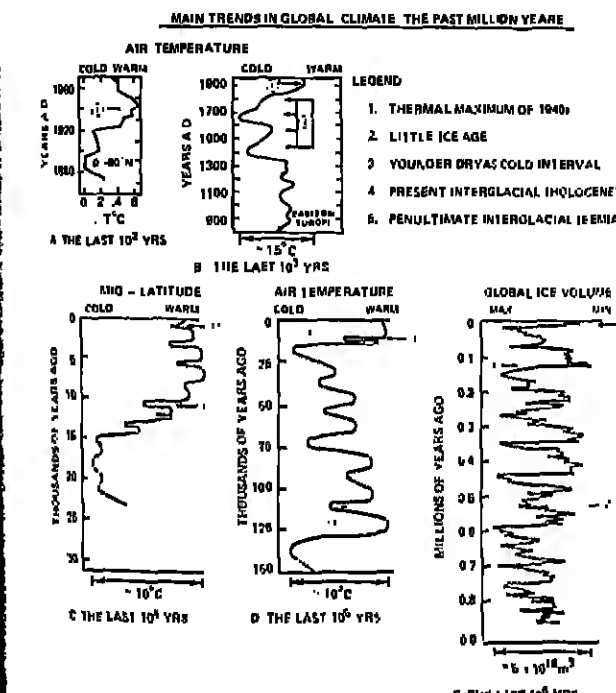


Fig. 3. Mean trends in global climate over the past 1 million years. (From Report of the Ad Hoc Panel on Present Interglacial, Federal Council for Science and Technology, 1974.)

Although one of the most widely quoted climatic curves, it is one of the most perplexing to explain. This temperature record is characterized by large annual changes which tend to obscure trends in the curve. Temperatures have declined since 1940 but have leveled off since 1965. Since then, surface temperatures have shown only a slight warming of 0.1° to 0.2°C. For average temperatures between surface and an elevation of 15,240 m, there has been no detectable change in temperature since 1885. Whether or not the fluctuations in this curve are natural or, in part, affected by anthropogenic factors is unknown. The curve has all of the principal characteristics of a temperature series produced by stochastic processes (Robock, 1979). The trend in temperature over this 100-year period is also inversely correlated with the transmissivity of the atmosphere (Bryson and Goodman, 1979). There exist many other climate time series for all or part of the last 100 years, including data on sea surface temperatures and atmospheric pressures. Barry et al. (1979) provide a short review of these climate indicators and what they say about short-term climate variability. Most of the data sets are relatively short (30 years or less) and cannot be used to document longer-term variability in this time period.

### Climate Variability: The Last 1000 Years

Figure 3b provides a general picture of climate variability over the past 1000 years. For this time interval there is no direct measure of climate comparable to the Northern Hemisphere temperature curve shown in Figure 3a. Rather, there are localized climate records compiled from observational, historical, and proxy data. Lamb (1988), for example, has compiled an index of winter severity in Europe from historical documents. LeMahieu (1974), using variations in tree ring widths as a proxy indicator of temperature and moisture, has reconstructed a near 1000-year climate record for mountain areas in California. Deserger et al. (1971) have developed a unique climate record for Greenland, based on isotopic chemical changes in ice cores; and Fritts et al. (1979) have reconstructed, from tree ring variations, a 400-year temperature, precipitation, and air pressure record for the United States. These and many other climate records indicate that the early part of the last millennium, from about 900 to 1200 A.D., was generally warm

and is referred to as the Medieval Warm Period. By contrast, the period between 1430 to 1850 was significantly cooler in Europe and eastern North America and is referred to as the Little Ice Age.

Some finer detail of climate variability over part of this time interval for the U.S. can be seen in the tree ring data analyzed by Fritts. Fritts' data allow a comparison of the general characteristics of each season for the past 400 years as reconstructed from tree ring variations. Some simple statistics can show how often severe winters like 1976-1977 and 1977-1978 have occurred in the past. During the 378 years from 1802 to 1978, the frequency of winters with a circulation pattern like 1976-1977 was 0.178 or 17.8 years per century. The frequency of winters like 1977-1978 was 8.6 years per century. Frequencies of winters like 1976-1977 varied the most from one century to another and were very frequent in certain time intervals. For example, the reconstructed circulation patterns between 1615 to 1685 resemble the winter of 1976-1977 with a frequency of 57.4 years per century. During the same time interval, the winters of 1977-1978 occurred only 12.5% of the time. From 1887 to 1929, no reconstructed winter circulation pattern resembled the winter of 1976-1977, and only 8% were similar to 1877-1978.

While there are many suggested causes for climate variability on this scale, a relationship with solar activity, as measured by sunspot occurrence, is often given prominence. While solar activity as measured by sunspot numbers has varied in a quasi-periodic fashion since the 1700's, there appears to have been a minimum of solar activity during the late 17th century. Eddy (1976), working from historic documents of visual sightings of sunspot activity, identified the period 1650-1710 as a low in sunspot activity. While Landsberg (1946) has recently identified, from newly studied historic diaries in Germany, a large number of sunspot and auroral observations made during the period 1685-1688, the total number of observed sunspots was still much less in the mid-17th century than at later periods. This period of time, termed the Maunder minimum, corresponds in time with a part of the 'Little Ice Age' in Europe. This correlation has been widely cited as a reliable link of sun and climate. It may not be so.

Historical data, by its very nature, is often incomplete and imprecise. Using such data as the sole basis for establishing a minimum of sunspot activity is therefore bound to be controversial. Reliable physical evidence that is accurately measured and global in representation does provide better proof of varying solar behavior. This evidence comes from carbon-14 fluctuations as observed and recorded in the annual growth of trees.

The production rate of carbon-14 in the upper atmosphere changes with time because the galactic cosmic ray flux responsible for C-14 production is modulated by changes in the magnetic properties of the solar wind. Changes in the atmospheric C-14 level are recorded in the annual growth of trees. Thus the C-14 levels derived from tree rings can be tied to the sun's modulation of the cosmic ray flux in the vicinity of the earth, and this provides a history of solar changes. Stuiver and Quay (1980) have determined the C-14 changes in trees over the past 1000 years. This C-14 record, used in conjunction with a carbon reservoir model that describes the terrestrial carbon exchange between the atmosphere, ocean, and biosphere, allows determination of a curve of changes in C-14 production rate (Figure 4a).

Because the C-14 production rates are dependent on neutron flux rates, which in turn are related to solar activity, the C-14 production rates should be compatible with and inversely related to sunspot activity. Stuiver and Quay have shown that the production rate index does correlate with observed sunspot behavior (Figure 4b, dashed line). From the C-14 production rates and the carbon reservoir model, Stuiver and Quay have been able to develop a theoretical long-term record of sunspot behavior. This proxy record (Figure 4b, solid line), which is fine tuned to the observed record, is characterized by two periods other than the Maunder minimum, when sunspot activity was low. The Spörer minimum occurs between 1418 and 1534, and the Wolf minimum between 1282 and 1342.

This important proxy record of sunspot behavior permits a direct test of possible links between solar minimum and climate. For example, do the times of the Spörer and Wolf minima coincide with periods of cooler climates? The results of such a comparison (Stuiver, 1980) indicate that there is no clear relationship, on this time scale, between sunspot behavior and climate. The earlier Wolf minimum appears to be coincident with a part of the Medieval warm period. It is thus becoming increasingly more difficult to link, in any straightforward fashion, sunspot and climate change.

### Climate Variability of the Last 10,000 and 1 Million Years

On these long time scales (Figure 3(c-e)), climate has been characterized by alternation between glacial and interglacial conditions. Over the past million years, ice ages have occurred many times, and only (on this time scale) that last 18,000 years ago a large part of the Northern Hemisphere lay under thousands of feet of ice. The last 10,000 years have been characterized by processes leading to deglaciation and subsequent evolution of modern climate. Because of the very important and exciting work by Hays et al. (1977), we now know that a major factor in the timing of past ice ages over the last million years has been due to changes in solar radiation received by the earth as a result of changes in the geometry of the earth's orbit. These orbital changes move and tilt the earth away from and toward the sun with a frequency of 19,000, 23,000, 41,000, and the sun with a frequency of 18,000 years has elapsed since the last ice age, a model of future climates, based on orbital theory and ignoring anthropogenic effect, predicts that the long-term trend over the next several thousand

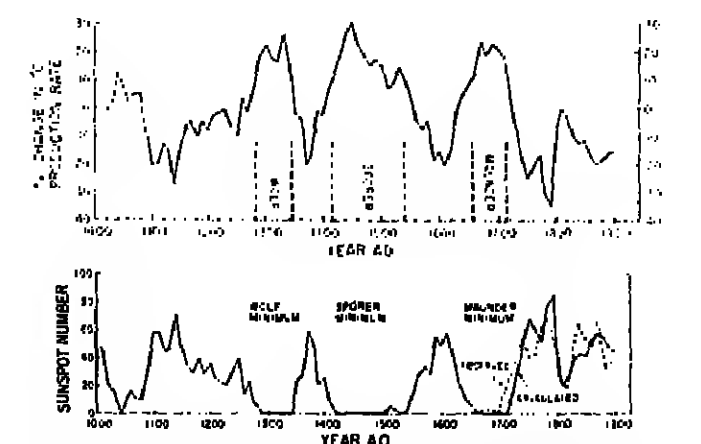


Fig. 4. (a) Changes in C-14 production rate calculated from carbon reservoir model relative to the average 1000 to 1860 production level. (b) Sunspot numbers as observed (dashed line) and calculated from production rate (solid line). [After Stuiver and Quay (1980) and reproduced with permission of the authors.]

years is toward glacial conditions. It is against this long-term trend that anthropogenic factors must also be measured.

While there are exciting things to say about climate variability on these long time scales, it is beyond the scope of this discussion, which emphasizes shorter-term climate variability. Long-term climate change is nicely discussed in Hecht (1979), Barry et al. (1980), and Mitchell (1970), who also provides an elegant discussion of climate variability in general.

### CO<sub>2</sub> Effect on Climate

Long-term future changes in the earth's climate may be related to the burning of fossil fuels. This comes about because the burning of these fuels releases large amounts of carbon dioxide (CO<sub>2</sub>) into the atmosphere. CO<sub>2</sub> is a gas which absorbs infrared radiation emitted by the earth's surface, and thus as its concentration in the atmosphere increases, so does the amount of heat it traps on the earth's surface. This 'greenhouse' effect may result in a global warming of a magnitude exceeding anything seen on the earth for millions of years.

It is not a recent hypothesis that man is affecting his environment by increasing the concentration of CO<sub>2</sub> in the atmosphere. As early as 1938, G. Callendar recognized that man, through the burning of fossil fuels, could change the composition of the atmosphere and affect climate. Nearly 20 years later, Revelle and Suess (1957) put the CO<sub>2</sub> question into global perspective. They said:

Human beings are now carrying out a large scale geophysical experiment of a kind which could not have happened in the past nor be repeated in the future. Within a few centuries we are returning to the air and oceans the concentrated organic carbon stored over hundreds of millions of years.

It is now nearly 23 years later, and in 1980 the documentation for the rise of CO<sub>2</sub> in the earth's atmosphere is at hand. The proof comes from direct measurements of CO<sub>2</sub> in the atmosphere at Mauna Loa, Hawaii, and other monitoring stations.

In 1957, as part of a research program developed for the

### Formulating A National Materials Policy: Public and Private Sector Roles

A conference to be held by the Department of Engineering and Public Policy at Carnegie-Mellon University, Pittsburgh, Pennsylvania, March 24, 1982

- Program Summary:
- The Need for a Federal Materials Policy: Competition with other Policies. Joel S. Hirschhorn, Project Director, Office of Technology Assessment.
  - The Role of Congress in Developing a National Materials Policy. Doug Weigert, Chairman, House Subcommittee on Science, Research and Technology, U.S. Congress.
  - The Aluminum Experience with Stockpiles. Charles W. Parry, President, Aluminum Company of America.
  - Materials Education in Relation to National Policy Making. Morris Cohen, Professor Emeritus, Massachusetts Institute of Technology.
  - Some Industrial Views on National Materials Policy. Julius J. Harwood, Director, Materials Sciences Laboratory, Ford Motor Company.
  - The Role of Economic Analysis. Leonard L. Fischman, President, Economic Associates Inc.
  - Can Materials Policy Become a Part of National Policy? Lindsey D. Norman, Vice President-Research, J. I. Steel.

Pre-registration by March 5, 1982 is recommended. Persons who do not pre-register should contact Paul Wynnblatt (412) 578-8711 before attending. Registration fees for this conference are:

Pre-registered by March 5: \$100  
On site registration: \$120  
Students: \$25

For more information please contact: Dr. Paul Wynnblatt, Department of Engineering and Public Policy, Carnegie-Mellon University, Pittsburgh, PA 15213, (412) 578-8711.

### Nuclear Regulatory Commission

Battelle Pacific Northwest Laboratory

## Partially Saturated Flow and Transport A Symposium

Of primary concern in the safe disposal of wastes is the environmental effect of near-surface disposal. Therefore, the Nuclear Regulatory Commission, in conjunction with Pacific Northwest Laboratory, is sponsoring a symposium to evaluate the current technology of flow and transport modeling in the partially saturated zone. The symposium on partially saturated flow and transport will emphasize recent work in both areas and will identify existing and future problems related to partially saturated flow and transport.

The technical program will cover two things and will include such topics as:

- Consolidation of Partially Saturated Soils
- Deterministic and Stochastic Models for Transport
- Parameters Governing Flow and Transport

Invited speakers from private industry, universities, and government agencies will present papers and open discussion sessions will be held.

The symposium will be held at the Battelle-Seattle Conference Center in Seattle, Washington, March 23-24, 1982. The Center, which is only ten minutes from downtown Seattle, provides a retreat-type environment with easy access to airport and other transportation facilities. For further information, contact:

Lorne Slominski  
Battelle Seminars and Studies Program  
4800 NE 41st Street  
P.O. Box C-5395  
Seattle, WA 98105  
(206) 527-5588

For registration information, contact Lorne Slominski no later than January 15, 1982. The registration fee is \$75.00 and is required by March 1, 1982. Details and registration forms are available for inquiries received through January 15. Attendance is limited. Registration will be on a first come, first serve basis. Advance registration is mandatory.



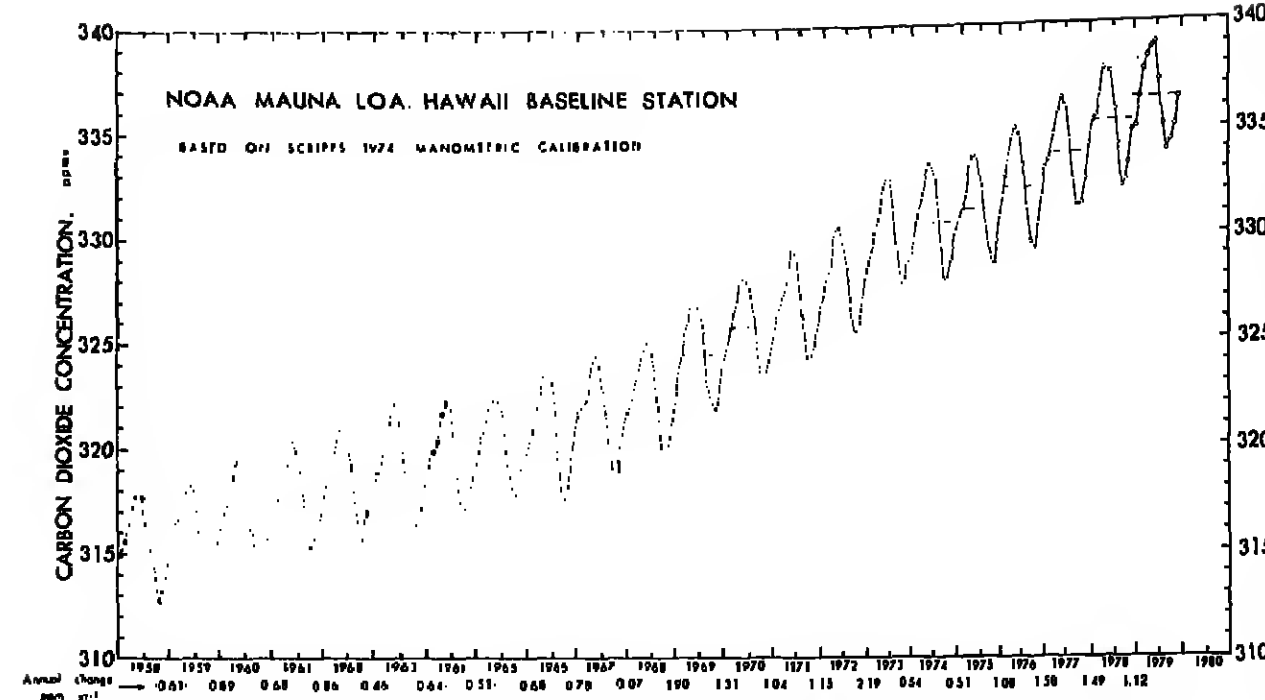


Fig. 5. CO<sub>2</sub> level recorded at Mauna Loa, Hawaii (in ppm) and annual changes. (Reproduced courtesy of NOAA.)

International Geophysical Year, laboratories were established at Mauna Loa (elevation 3400 m) and the South Pole to begin accurate and regular measurements of CO<sub>2</sub> in the atmosphere. The monitoring at Mauna Loa has resulted in the observations shown in Figure 5. The curve clearly shows seasonal variations in respiration-photosynthesis, with an amplitude of about 6 ppm. Maximum CO<sub>2</sub> occurs in April, minimum in October. The decrease represents the excess photosynthesis uptake of CO<sub>2</sub> over decay and respiration during the Northern Hemisphere summer. In addition to the seasonal signal from these data it is clear that since 1958 the amount of CO<sub>2</sub> in the atmosphere has steadily increased. The current value of 336 ppm, or 700 × 10<sup>15</sup> g C, represents an increase of 20 ppm since measurements began in 1958. Estimates of the amount of CO<sub>2</sub> in the atmosphere prior to 1958 are between 265 and 290 ppm (550 to 620 × 10<sup>15</sup> g C). Thus between 80 and 150 × 10<sup>15</sup> g C have been contributed to the atmosphere since preindustrial days. CO<sub>2</sub> produced by industrial activity from 1880 to 1979 is about 160 × 10<sup>15</sup> g C (Table 2). Approximately 80 × 10<sup>15</sup> g C of this amount was contributed between 1958 and 1979. The source for these data on CO<sub>2</sub> emissions from the burning of fossil fuels are UN records, which may be subject to an error of 10% or 15%; the data are, however, continuous and internally consistent. From 1860 to 1970 the CO<sub>2</sub> emissions from fossil fuels grew at a rate of 4.3% per year, except for the periods of world wars. If CO<sub>2</sub> production continued at this rate, annual CO<sub>2</sub> production would approximately reach 14 × 10<sup>15</sup> g C by the year 2000 and 41.5 × 10<sup>15</sup> g C by 2025.

The increase in CO<sub>2</sub> production has declined over the past 10 years and is now about 3.6%. The amount of CO<sub>2</sub> emissions from fossil fuels can be projected reasonably for the next 20 years, since the time required to make major shifts in energy production or consumption is of this magnitude. Predictions beyond the year 2000 are much more difficult to make and are a product of complex interactions of demographic, economic, and social variables.

The measurements at Mauna Loa of CO<sub>2</sub> (Table 2) in the atmosphere over the period 1960–1979 show an increase of 9.5 ppm from about 317.2 to 336 ppm. This increase is 6‰ and is equivalent to an additional 39 × 10<sup>15</sup> g C.

Since the beginning of the CO<sub>2</sub> measurements at Mauna Loa, the observed increase has accounted for about 50% of the carbon dioxide released by the burning of fossil fuel and destruction of vegetation. The other 50% has been added to the other carbon reservoirs, which are the oceans and the biosphere. Estimates of CO<sub>2</sub> remaining in the atmosphere vary between 48% and 56% (Broecker, et al., 1980). While at first there was considerable discussion that the biosphere itself, through deforestation, was also a major contributor of CO<sub>2</sub>, it now appears that this contribution is small and that to a first approximation the fossil CO<sub>2</sub> released to the atmosphere can be adequately accounted for in existing carbon cycle models (Broecker et al., 1980). (This point is controversial, however, and I am treating it casually in this review, since the topic is mainly climate change. Further discussion is given in reviews of the global carbon cycle, for example Bolin et al. (1979).)

#### CO<sub>2</sub> Effect on Climate

The primary effect of an increase of CO<sub>2</sub> in the atmosphere is to cause more atmospheric absorption of thermal radiation from the earth's surface and thus to increase air temperature. Numerical modeling of this process with global atmospheric general circulation models (GCM) suggest a global warming of the earth of about 2°C with a doubling of CO<sub>2</sub> and 4°C with a quadrupling. The model experiments indicate that the warming is greatest in polar regions, where the increase may be 3 times as large as in tropical regions. Climate simulations with increased levels of CO<sub>2</sub> also provide estimates of changes in the pattern of evaporation and precipitation and in the extent of sea ice. The value of these experiments is primarily as diagnostic tests of climate models and their intercomparison. Although present climate models do capture the main large-scale features of the atmosphere, they are severely limited in portraying oceanic-atmospheric interactions, cloudiness, and detailed regional climate changes. Large-scale ocean modeling comparable to existing atmospheric modeling is not

presently available because of both lack of observational data on appropriate synoptic scales and inadequate understanding of major oceanic mixing and circulation processes. Thus present GCM's of both atmosphere and ocean are capable only in a modest way of duplicating the observed climate. Simulations of climate with increased levels of CO<sub>2</sub> must be viewed in the context of the capabilities of these models to simulate modern climate. (A review of the strengths and weaknesses of climate modeling is beyond the scope of this report, but I recommend the referenced papers given by Barry et al. (1979) in their review of climate change.) Herein, I can discuss only the most recent results of large-scale climate modeling with increased CO<sub>2</sub> levels and compare the results to previously published reports.

Manabe and Stouffer (1980) simulate global climate with 2 and 4 times the present level of CO<sub>2</sub>, using the sophisticated general atmospheric circulation and simplistic ocean model developed at the Geophysical Fluid Dynamics Laboratory (GFDL). The model consists of an atmospheric GCM and a mixed-layer ocean model with uniform thickness. Like most GCM's this model predicts changes in vorticity components of vorticity, divergence, temperature, moisture, and surface pressure from the basic equations of motion, thermodynamics, and continuity.

The ocean model is a static isotherm water layer of uniform 68 m thickness. This thickness assures that the heat storage associated with the annual cycle of sea surface temperature is included in the model.

The model is run beginning with isothermal, dry, and motionless atmosphere and with an atmospheric concentration of CO<sub>2</sub> at 300 ppm. Stable climate conditions develop after about 10 years of model time. The control experiments successfully reproduce the observed basic features in the seasonal variation of the atmosphere. In response to a quadrupling of the CO<sub>2</sub> level of the atmosphere, the model produces a new equilibrium climate which shows an overall global average increase of 4.1°C in surface temperatures. Low-latitude changes are on the order of 3–4°C; high-latitude changes are 6–8°C in the Southern Hemisphere and 8–14°C in the Northern Hemisphere. Figure 6 shows the latitude height distribution of the difference in zonal mean air temperatures between an atmosphere with the present and 4 times the amount of CO<sub>2</sub>. Estimated temperature changes are half as great for a doubling of CO<sub>2</sub> levels.

Manabe and Stouffer (1980) give an excellent discussion of the results of their experiment with regard to the latitudinal and seasonal variation of the changes in precipitation, evaporation, and sea-ice distribution. In general the model shows greater moisture content of the air and an increase in the poleward transport of moisture. Additional moisture generated in the tropics is transported to high latitudes, and both precipitation and runoff rate increases. As temperatures increase in the Northern Hemisphere, sea ice is reduced. With 4 times the CO<sub>2</sub> level in the atmosphere, sea ice disappears completely from the Arctic Ocean during a few summer months.

The global model used in this study contains many simplifications and idealizations. Some important physical processes, such as the horizontal heat transport of ocean-

ocean currents are not considered. In attempting to simulate the present climate, the surface air temperature over the entire circum-Antarctic Ocean is overestimated, resulting in the underestimation of the area covered by sea ice.

The results of this model suggest a somewhat lower global temperature increase than previously estimated by these and other authors. The differences are not great, and there is a general convergence of a ±2°C temperature increase for a doubling of CO<sub>2</sub>. This number is generally higher than estimates derived from simple radiation balance models, which for the most part record only the expected atmospheric response to CO<sub>2</sub> increase in the atmosphere independent of atmospheric and oceanic feedback processes. For example, Newell and Dopplick (1979) assume that the CO<sub>2</sub>-induced change of temperatures and mixing ratio of water of surface air is zero when they evaluate the CO<sub>2</sub>-induced changes in sensitive and latent heat flux from the earth's surface to the atmosphere. Thus the warming of surface temperature is greatly underestimated.

Most of these models suggest a greater warming in the polar regions than in the tropical ones. Since the West Antarctic ice sheet is thought to be relatively unstable in comparison to the remainder of the ice cover over Antarctica, there is concern that this ice sheet might disintegrate or surge because of the temperature increase. There is, however, considerable disagreement among glaciologists about the likelihood of a collapse of the West Antarctic ice sheet. A recent meeting of experts (Orono, Maine 1980), sponsored by the Department of Energy, produced recommendations for a research program to clarify conflicting opinions.

It is not clear at this time how to verify that any global increase in temperature (should it occur over the next decades) is due to CO<sub>2</sub>. Because the intermediate layers of the ocean are expected to absorb some of the increased heat, any atmospheric increase in temperatures may be delayed behind the CO<sub>2</sub> input by perhaps several decades (National Academy of Sciences, 1979). Thus it is not obvious how a global warming, such as that which occurred between 1850 and 1940, presumably due to non-CO<sub>2</sub> effects, may be distinguished from a predicted warming due to CO<sub>2</sub>.

If average global temperatures were indeed to increase, new patterns of evaporation and precipitation would likely develop. The effect of such a change would be felt everywhere. The Manabe-Stouffer experiment discussed above suggests that some regions would become wetter, others drier, most warmer, and some colder. The ultimate consequence would be a global society and a global ecosystem which would be forced to adapt to a new climatic state with a different distribution of temperatures and precipitation, winds, humidity, and the like. How climate variability would change as a result of changes in CO<sub>2</sub> level is unknown. It is, however, variability of climate, more than slow climate change, which affects the economic and social well being of society.

#### Living with Climate Change

As a theme for this article, I have centered on drought as an example of a climatic extreme that has significant impact on society. While a drought of the magnitude of the 1930's has not occurred since, other climatic variations from drought to extreme cold have been characteristic of the past decade. As discussed above, the perception that climate is becoming more variable has given rise to international and national programs designed to understand better the causes of climatic change and to utilize better existing knowledge of climate variability in decision making and resource management.

The impact of climate on society is both a product of the climate change itself and the vulnerability of society. Whether society today is more or less vulnerable to major climatic changes (than in the past) is a research question for the decades ahead. Even given no climate change, can society manage its affairs with increased population, energy, and food demands. The report of the Council on Environmental Quality, Global 2000, suggests a grim future picture for world society as the result of overpopulation, limited fuel resources, and severe water shortages. Water availability may, in fact, be the single most important environmental variable in the decades ahead.

For the past 30–40 years, the normal water supply in most U.S. river basins has been adequate for agricultural, industrial, and municipal purposes. As population increases and industry develops (particularly in the Southwest), the balance between available water and water needs be-

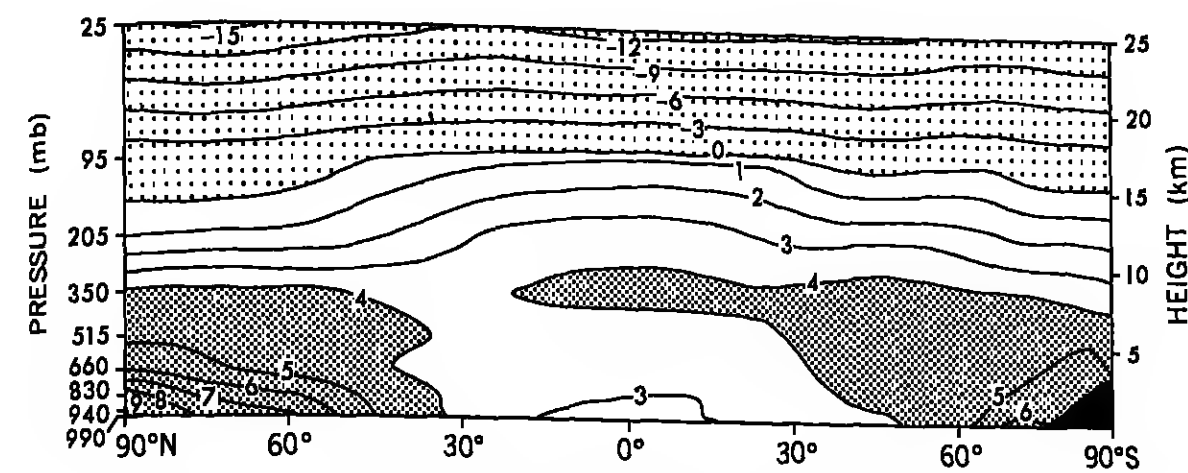


Fig. 6. Zonal mean difference in annual mean temperature (degrees Kelvin) of the model atmosphere between 4 times CO<sub>2</sub> and present levels. (From Manabe and Stouffer and reproduced with permission of the authors.)

comes critical. For part of the Colorado River Basin, the water shortage is already quite evident.

Suppose there is a climate change of some degree between now and the year 2000, what would be the effect on the 18 major water regions of the U.S.? Stockton and Boggess (1979) have made a preliminary comparison of present supply and demand for water, with projected values for a scenario of ±2° warmer or colder and ±10% greater or less precipitation. In general, most regions east of the Rocky Mountains would not be drastically influenced by the hypothetical climatic changes above. Local problems of flooding, transportation, or waste management could be met by alternate management strategies.

River basins west of the Mississippi River, however, could experience significant shortages under a warmer and drier scenario. Stockton has calculated, using his hydrologic model, that the increased water evaporation from water surfaces, soil, and plants caused by a rise of 2°C in mean annual temperature accompanied by a 10% decrease in total precipitation could result in decreases of 40% to 60% in annual surface water supplies. Climate changes of this magnitude have occurred naturally over the past 150 years. The regions that would suffer major impacts would be Arkansas, White-River, Texas-Gulf, Rio Grande, Upper and Lower Colorado, California, and Missouri. As groundwater reserves are already heavily utilized in these seven regions, it cannot be considered a viable alternative supply to supplement surface water shortages.

For the climatic scenario of cooler (by 2°) and wetter (by 10%), the national impact would be mostly beneficial. Regions that would be mildly adversely affected because of excess water would include the South Atlantic Gulf, New England, Lower Mississippi, and Great basins.

Thus the danger lies ahead for the western U.S., where, under drier conditions, severe water shortages can be expected. Even in the absence of any climatic changes, water shortages may be likely because of the increased need for water in the development of energy sources, for agriculture, and for the increased industrialization and expansion of municipal areas. While planning for water access has been done for many years (flood control, zoning, land management, etc.), planning for water shortages is not well advanced. Considering that most groundwater sources in the western U.S. are being used up faster than they are replenished, the problem of water management in the western U.S. may be one of the most serious problems of the year 2000.

In fact, it may be the problem of water availability that determines how society may respond to CO<sub>2</sub> climate change. A report of the National Academy of Sciences on how CO<sub>2</sub> induced climate changes might affect society concluded that

... changes in availability of water are the single most significant consequence of climate change through the next century—while modest precipitation increases in areas well supplied at present could be accommodated, similar decreases in some currently marginal arid and semi-arid regions and increases in the frequency of drought could have serious impacts.

Food and water are intimately related, and I conclude this long essay by again returning to drought and agriculture in the Midwest. This feat of growing corn in a semiarid region like western Kansas has been made possible by heavy irrigation of the groundwater from the Pliocene age Ogallala formation, which underlies parts of the high plains. In most of the high-plain region, groundwater withdrawals are far in excess of recharge. To meet the demands of agriculture and population in this area in the year 2000 will require extensive water management systems, such as the existing groundwater management districts (GWM), which permit users to determine the level of water consumption. Additional options for supplementary groundwater demands may involve the transfer of water from the Missouri River or other water basins. Such projects would involve gargantuan costs—even by today's monetary standard; or in the extreme case, with diminishing water resources, western Kansas, like areas of Texas, could revert to overgrazing. More likely, as reported by John Welsh (What To Do When the Well Runs Dry, Science, 210, 754–758, 1980), western Kansas could change from irrigated corn agriculture to the raising of less water-intensive crops and, perhaps ultimately, to dryland farming of wheat and grain sorghum.

The problems for U.S. farmers today, like farmers during the 1930's and like Indiana of Mill Creek, is living and working in a world with a climate that is unpredictable from year to year. Unlike the Indians of Mill Creek we have extensive technology available to us to insulate society from extreme weather or climate events. Unlike the Indians of Mill Creek, abandonment of the land is not our only option. But like the

Indians of Mill Creek we remain strongly affected by climate. It is one natural resource that is still a challenge to men.

#### Acknowledgments

I am grateful to several colleagues who reviewed drafts of this article and who corrected many of my silly mistakes. They are: Lester Mehta and William Elliott (NOAA), who also provided Figure 6; Reid Bryson (U. Wisconsin), John Perry (National Academy of Sciences); Richard Warrick (NCAR), who also gave permission to reproduce Figures 1 and 2; Ken Bergman (NSF); and J. Murray Mitchell (NOAA). Figure 7 was also provided by S. Manabe (GFDL); Figure 5 by Minze Stouffer (U. Washington). Uncorrected mistakes are my own, and opinions expressed in this article are mine and do not represent the official position of the National Science Foundation.

#### References

- Barry, R. A., D. Hecht, J. Kutzbach, W. D. Sellers, T. Webb, III, and P. B. Wright, Climatic change, *Rev. Geophys. Space Phys.*, 17, 1603–1612, 1979.
- Bolin, B., E. T. Degens, S. Kempe, and P. Keiner (Eds.), *The Global Carbon Cycle, Scope 13 Report*, 491 pp., John Wiley, N.Y., 1979.
- Broecker, W. S., T. Takahashi, H. J. Simpson, and T. H. Peng, Fate of fossil fuel CO<sub>2</sub> and the global carbon budget, *Science*, 207, 1041–1044, 1980.
- Bryson, R. A., and D. Bernier, Climatic change and the Mill Creek culture of Iowa, Part 1. Chapt. 1, Introduction and Project Summary, *J. Iowa Archaeol. Soc.*, 15, 1–34, 1966.
- Bryson, R., and B. M. Goodman, Volcanic activity and climatic changes, *Science*, 1041–1044, 1980.
- Bryson, R. A., D. Bernier, and W. M. Wendland, The character of late glacial and post glacial climatic changes, Pleistocene and Recent Environments of the Central Great Plains, *Spec. Publ. 3*, pp. 53–74, Univ. of Kansas, 1970.
- Cellender, G., On the amount of CO<sub>2</sub> in the atmosphere, *Tellus*, 10, 243–246, 1958.
- Cheney, J. C., Dynamics of desert and drought in the Sahel, *Q. J. R. Meteorol. Soc.*, 101, 193–202, 1975.
- Chico, T., and W. D. Sellers, Interannual temperature variability in the United States since 1866, *Clim. Change*, 2, 139–148, 1979.
- Dansgaard, W. S., S. J. Johnson, H. B. Clausen, and C. C. Langway, Jr., Climatic record revealed by the Camp Century ice core, In: *The Late Cenozoic Glacial Ages*, edited by K. Turekian, pp. 37–56, Yale University Press, New Haven, 1971.
- Eddy, J. A., *The Medieval minimum*, *Science*, 192, 1169–1202, 1976.
- Fritts, H. C., *Tree Rings and Climate*, Academic, New York, 1976.
- Fritts, H. C., R. Lohgen, and G. A. Gordon, Variations in climate since 1602 as reconstructed from tree rings, *Q. Res.*, 12, 18–46, 1979.
- Glantz, M. H., Nine lectures of natural disaster: The case of the Sahel, *Clim. Change*, 1, 69–84, 1977.
- Hayes, D. J., J. Imbrie, and N. J. Shackleton, Variations in the earth's orbit: Pacemaker of the ice Age, *Science*, 194, 1121–1132, 1977.
- Hecht, J. D. (Ed.), Paleoclimatic research: Status and opportunities, *Quat. Res.*, 10, 6–17, 1979.
- LaMerche, V. C., Jr., Paleoclimatic inferences from long tree-ring records, *Science*, 183, 1043–1046, 1974.
- Lamb, H. H., Climatic fluctuations, In: *World Survey of Climatology*, 2, General Climatology, edited by H. Flohn, pp. 173–249, Elsevier, New York, 1969.
- Landsberg, H., Climate as a natural resource, *Sci. Mon.*, 63, 293–296, 1946.
- Ludlum, D. M., *Weather Record Book*, Weatherwise, Inc., Princeton, N.J., 1971.
- Manabe, S., and R. L. Stouffer, Sensitivity of a global climate model to an increase of CO<sub>2</sub> concentration in the atmosphere, *J. Geophys. Res.*, 85, 5526–5554, 1980.
- McQuigg, J. P., L. Thompson, S. LeDuc, M. Locard, and G. McKey, The influence of weather and climate on U.S. grain yields: Bumper crops or drought, report, NOAA, U.S. Dep. of Commerce, Washington, D.C., 1973.
- Mitchell, J. M., Jr., An overview of climatic variability and its causal mechanisms, *Q. Res.*, 2, 461–494, 1976.
- Mitchell, J. M., Jr., C. W. Stockton, and D. M. Meko, Evidence of a 22-year rhythm of drought in the 17th century, In: *Solar-Terrestrial Influences on Weather and Climate*, edited by B. M. McCormac and T. A. Seliga, pp. 125–143, D. Reidel, Dordrecht, Holland, 1979.
- National Academy of Sciences, CO<sub>2</sub> and Climate, A Scientific Assessment, report, Nat. Acad. Sci., 22 pp., Washington, D.C., 1979.
- Newell, R. E., and T. G. Dopplick, Questions concerning the possible influence of global CO<sub>2</sub> on atmospheric temperatures, *J. Appl. Meteorol.*, 18, 622–625, 1979.
- Newman, J. E., Drought impacts on American agricultural productivity, In: *North American Drought*, edited by N. J. Rosenberg, pp. 43–82, Westview Press, Boulder, Colo., 1978.
- Palmer, W. C., Meteorological drought, *Res. Pap.*, 45, 58 pp., U. S. Dep. of Commerce, Washington, D.C., 1965.

- Ravett, R., and H. Seuss, CO<sub>2</sub> exchange between the atmospheric CO<sub>2</sub> during the past decade, *Tellus*, 9, 18–27, 1957.
- Robock, A., Internal and external causes of climate change, *J. Atmos. Sci.*, 35, 1111–1122, 1978.
- Stockton, C. W., and W. R. Boggess, Geohydrological Implications of Climate Change on Water Resource Development, report, 206 pp., U.S. Army Coastal Eng. Res. Center, Fort Belvoir, Va., 1979.
- Stouffer, M., Solar variability and climatic change during the current millennium, *Nature*, 286, 868–871, 1980.
- Stouffer, M., and P. D. Quoy, Changes in atmospheric carbon-14 attributed to a variable sun, *Science*, 207, 11–19, 1980.
- Warrick, R., Drought in the Great Plains: A case study of research on climate and society in the U.S., in *Climatic Constraints and Human Activity*, edited by J. Ausubel and A. K. Blawie, pp. 63–124, NASA Proc. Ser., Pergamon, N.Y., 1980.
- Warrick, R., and R. Kales, Testing hypotheses about the effects of climate fluctuations on society: Three case studies, paper presented at Annual Meeting AAAS, Toronto, January 1981.



Alan Hocht is director of the Climate Dynamics Program, Division of Atmospheric Sciences, National Science Foundation. He is a fellow of the Geological Society of America, president of INQUA paleoclimatology commission, a member of U.S. National INQUA committee, associate editor of *Climatic Change*, and chairman of the American Meteorological Society's Committee on Climatic Variations. While trained as a geologist, he has broad interest in modern and past climatic variations, climate modeling, and the impacts of climate on society.

## News

### COSOD Update

The Conference on Scientific Ocean Drilling (COSOD) (EOS, December 1, p. 1152) identified a set of global scientific objectives ranging from the continental margins to the deep sea that will require a worldwide program of drilling in the Atlantic, Pacific, Indian, and polar oceans, explained Roger L. Larson, chairman of the COSOD Steering Committee.

However, COSOD did not aim to provide scientific goals for the Ocean Margin Drilling Program (OMDP). The main

The 12 top-priority scientific programs, with relevant questions, identified at COSOD are listed below in nonpreferential order. This list is still subject to revision by the COSOD Steering Committee and will almost certainly evolve as the future ocean drilling program proceeds.

Processes of magma generation and crustal construction at mid-ocean ridges. What is the composition of the oceanic layer? Is the ophiolite analogy a valid model for ocean crust?

The configuration, chemistry, and dynamics of hydrothermal systems. What are the dimensions and characteristics of open versus closed and active versus passive hydrothermal systems?

Early rifting history of passive continental margins. What is the shallow and deep structure of attached and detrital-lifted margins versus those characterized by extensive volcanism?

The dynamics of tectonic evolution. What are the relative motion, deformation, and pore-water characteristics of sediments being subducted at accreting and nonaccreting margins?

Forearc to back-arc structure and magmatic history. What are the space and time relationships of back-arc spreading, compression, and volcanism at island arcs? The response of marine sedimentation to fluctuations in sea level.

Which on-land, off-land sequences and intervening unconformities represent sea-level fluctuations and which represent vertical tectonic motion? What is the response of deep-sea sedimentation to sea-level fluctuations?

Sedimentation in oxygen deficient oceans. What are the ocean circulation, paleoclimatic, and potential hydrocarbon characteristics associated with Cretaceous black-shale deposits?

Global mass balancing of sediments. What are the best estimates of the world sediment mass and composition balances in space and time?

Ocean circulation history. What is the response of ocean circulation to changing boundaries, especially the Drake Passage, the Isthmus of Panama, and the Tethys Seaway? How does this vary in glacial and nonglacial eras?

The response of the atmosphere and oceans to orbital variations. How have gravitational interrelations with other planets, especially Jupiter, affected paleocirculation in the atmosphere and hydrocycle?

The history of the earth's magnetic field. What is the nature of the magnetic field during a magnetic reversal? What is the detailed reversal and paleointensity history of the magnetic field in the past 200 million years?

The process and mechanism of evolution in marine organisms. Have the details of evolution been characterized by continuous change or by punctuated equilibrium states in marine organisms?



